

$$\begin{aligned} \operatorname{curl} \mathbf{u} &= \hat{\mathbf{e}}_r \frac{1}{r \sin \theta} \left( \frac{\partial}{\partial \theta} (\sin \theta u_\phi) - \frac{\partial u_\theta}{\partial \phi} \right) \\ &+ \hat{\mathbf{e}}_\theta \frac{1}{r \sin \theta} \left( \frac{\partial u_r}{\partial \phi} - \sin \theta \frac{\partial}{\partial r} (ru_\phi) \right) \\ &+ \hat{\mathbf{e}}_\phi \frac{1}{r} \left( \frac{\partial}{\partial r} (ru_\theta) - \frac{\partial u_r}{\partial \theta} \right) \end{aligned} \quad (16)$$

$$\begin{aligned} \nabla^2 \psi &= \frac{1}{r^2} \frac{\partial}{\partial r} \left( r^2 \frac{\partial \psi}{\partial r} \right) + \frac{1}{r^2 \sin \theta} \frac{\partial}{\partial \theta} \left( \sin \theta \frac{\partial \psi}{\partial \theta} \right) \\ &+ \frac{1}{r^2 \sin^2 \theta} \frac{\partial^2 \psi}{\partial \phi^2}. \end{aligned} \quad (17)$$

These expressions are used when we discuss spherical waves in Section 2.4 and the earth's normal modes in Section 2.9.

A final point worth noting is that the elements of volume and surface used in integrals are different in spherical coordinates from rectangular coordinates. In spherical coordinates (Fig. A.7-5) there are several scale factors, so an element of surface area is

$$dS = r^2 \sin \theta d\theta d\phi, \quad (18)$$

and an element of volume is

$$dV = r^2 \sin \theta dr d\theta d\phi. \quad (19)$$

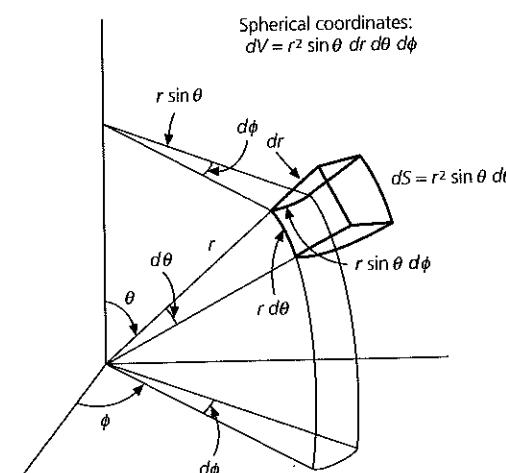


Fig. A.7-5 Definition of the element of volume in spherical coordinates. Unlike the case of Cartesian coordinates, the volume element in spherical coordinates is not a cube. (Marion, 1970. From *Classical Dynamics of Particles and Systems*, 2nd edn, copyright 1970 by Academic Press, reproduced by permission of the publisher.)

## A.8 Scientific programming

Most seismological applications require computers, and these requirements, especially in exploration applications with very large data volumes, have spurred the development of computer software and hardware. Some remarks about the use of computers in seismology thus seem appropriate.

Computer usage in seismology includes several broad and overlapping categories:

- Computers are often used in data acquisition and recording systems.
- Data are initially displayed and manipulated using computers.
- Subsequent analysis is frequently done using computers. For example, seismograms can be filtered to enhance certain frequencies or combined to better resolve certain features.
- Theoretical, or *synthetic*, seismograms are often computed for a range of the parameters under study and compared to data to find the best fit.
- Computers are used to *invert* seismological data to determine the parameters of a model which best matches the data.
- Computer modeling is often used to draw geological inferences from seismological observations. For example, seismic velocity data are compared to the predictions of models for the velocity of rock as a function of composition, temperature, and pressure.

These applications often require *scientific programming*, a programming style used for essentially mathematical applications. Some problems in this book also require scientific programming. Although programming is a matter of personal style, this section discusses several points that may be helpful. The suggested reading provides some starting points for readers interested in pursuing these topics further.

### A.8.1 Example: synthetic seismogram calculation

Consider a program to compute a synthetic seismogram for waves in a one-dimensional constant-velocity medium. A mathematically idealized string that illustrates features of wave behavior. The program is based on  $u(x, t)$ , the displacement as a function of position  $x$  and time  $t$ . The displacement is zero at the fixed ends of the string,  $x = 0$  and  $x = L$ , between which waves travel at speed  $v$ . As in Section 2.2.5, the displacement can be written as the sum of the normal modes of the string, each of which is a standing wave with  $n$  half wavelengths along the string,

$$u_n(x, t) = \sin(n\pi x/L) \cos(\omega_n t),$$

and vibrates at a characteristic frequency, or eigenfrequency,  $\omega_n = n\pi v/L$ .

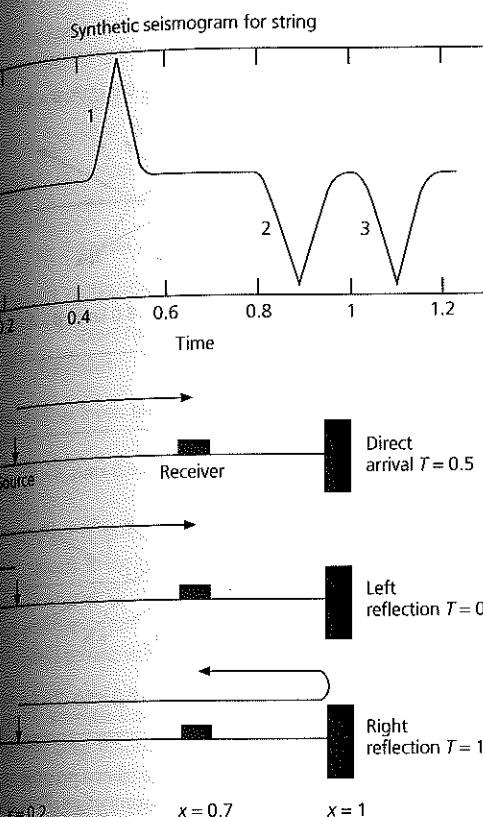


Fig. A.8-1 Synthetic seismogram for a string showing the direct wave (1) and reflections (2, 3) from both ends. Bottom: Geometry of source and receiver positions, and the times of the direct and arrivals.

at position  $x_s$  generates a pulse at time zero with amplitude  $a$ . The propagating waves are described by a weighted sum of modes

$$u(x, t) = a \sin(n\pi x/L) \sin(n\pi x_s/L) \cos(\omega_n t) \exp[-(\omega_n t)^2/4]. \quad (3)$$

The displacement  $u(x, t)$  for any position and time, called a "stringogram" giving the displacement versus a receiver position  $x_r$  is  $u(x_r, t)$ . Alternatively, a plot of the displacement everywhere on the string at time  $t$ ,

is a program to evaluate a synthetic seismogram for simplicity, we use a string of length 1 m<sup>1</sup> and speed 1 m/s, a source at  $x_s = 0.2$  m and a receiver at  $x_r = 0.7$  m. To approximate the infinite sum, the program adds terms. The seismogram (Fig. A.8-1, top) is calculated over 50 time steps covering 1.25 s. This program is written in

Fortran, a language that is especially suitable for scientific programming and is therefore commonly used in seismology (and thus in this book). The program could be also written in other languages, but the general points would still apply.

```
C SYNTHETIC SEISMOGRAM FOR HOMOGENEOUS STRING
C DISPLACEMENT U AS FUNCTION OF TIME T
C CALCULATED BY NORMAL MODE SUMMATION
DIMENSION U(200)
PI = 3.1415927
C
C PARAMETERS (NORMALLY WOULD COME FROM INPUT)
C STRING LENGTH (M)
ALNGTH = 1.0
C VELOCITY (M/S)
C = 1.0
C NUMBER OF MODES
NMODE = 200
C SOURCE POSITION (M)
XSRC = 0.2
C RECEIVER POSITION (M)
XRCVR = 0.7
C SEISMOGRAM TIME DURATION (S)
TDURAT = 1.25
C NUMBER TIME STEPS
/ NTSTEP = 50
C TIME STEP (S)
DT = TDURAT/NTSTEP
C SOURCE SHAPE TERM
TAU = .02
C
C LIST PARAMETERS
WRITE (6,3000)
3000 FORMAT('SYNTHETIC SEISMOGRAM FOR STRING')
WRITE (6,3001) NMODE
3001 FORMAT('NUMBER OF MODES', I6)
WRITE (6,3002) ALNGTH, C
3002 FORMAT ('LENGTH (M)' F7.3, 'VELOCITY,
X (M/S)', F7.3)
WRITE (6,3003) XSRC, XRCVR
3003 FORMAT ('POSITION (M): SOURCE', F7.3,
X 'RECEIVER', F7.3)
WRITE (6,3004) TDURAT, NTSTEP
3004 FORMAT ('SEISMOGRAM DURATION (S)', F7.3,
X I6, 'TIME STEPS')
WRITE (6,3005) TAU
3005 FORMAT ('SOURCE SHAPE TERM', F7.3)
C
C INITIALIZE DISPLACEMENT
DO 5 I = 1, NTSTEP
U(I) = 0.0
5 CONTINUE
C
C OUTER LOOP OVER MODES
DO 10 N = 1, NMODE
ANPIAL = N*PI/ALNGTH
```

